Modeling annual production and carbon fluxes of a large managed temperate forest using forest inventories, satellite data and field measurements

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Summary We evaluated annual productivity and carbon fluxes over the Fontainebleau forest, a large heterogeneous forest region of 17,000 ha, in terms of species composition, canopy structure, stand age, soil type and water and mineral resources. The model is a physiological process-based forest ecosystem model coupled with an allocation model and a soil model. The simulations were done stand by stand, i.e., 2992 forest management units of simulation. Some input parameters that are spatially variable and to which the model is sensitive were calculated for each stand from forest inventory attributes, a network of 8800 soil pits, satellite data and field measurements. These parameters are: (1) vegetation attributes: species, age, height, maximal leaf area index of the year, aboveground biomass and foliar nitrogen content; and (2) soil attributes: available soil water capacity, soil depth and soil carbon content. Main outputs of the simulations are wood production and carbon fluxes on a daily to yearly basis. Results showed that the forest is a carbon sink, with a net ecosystem exchange of 371 g C m–2 year –1. Net primary productivity is estimated at 630 g C m–2 year –1 over the entire forest. Reasonably good agreement was found between simulated trunk relative growth rate (2.74%) and regional production estimated from the National Forest Inventory (IFN) (2.52%), as well as between simulated and measured annual wood production at the forest scale (about 71,000 and 68,000 m3 year–1, respectively). Results are discussed species by species.

Keywords: CASTANEA process-based model, Fontainebleau forest, NEE, remote sensing, soil carbon, spatialization.

Introduction

There are two challenges in modeling forest production: coupling forestry knowledge and process-based models, and scaling up from the stand to the regional scale. Coupling the mechanistic approach based on ecophysiological rules with empirical forestry knowledge provides a means to improve the predictions of forest production under changing environmental conditions (Mäkelä et al. 2000). Early forest production models focused on forestry and yield prediction. These models used empirical rules based on large data sets from field experiments (Schober 1975, Dhote 1991) and could reproduce tree growth over a century, assuming no climatic trend, according to species, forest management and age of planting. They are not designed, however, to predict seasonal and inter-annual variations in tree growth and stand biomass increment and do not account for effects of global climatic change. Recently developed process-based forest ecosystem models couple water and carbon exchange between vegetation and atmosphere. Some of these models also consider litter and soil mineralization processes. These models are designed to predict soil organic matter dynamics and net ecosystem exchange of CO2 (NEE) (Baldocchi and Harley 1995, Hoffmann 1995, Kirschbaum 1999). Although most of these models were developed to study and quantify water and carbon fluxes, they are now used for productivity assessment on year to century time scales, in the context of climate change (He et al. 1999, Mäkelä et al. 2000, Landsberg 2003). These model simulations are in good agreement with actual measurements of carbon fluxes between the ecosystem and the atmosphere that have been measured in several forest stands by the eddy covariance technique (Running et al. 1999, Kramer et al. 2002, Law et al. 2002).

The coupling between forestry and ecophysiological models should first be done at the stand scale, because this is the scale at which the process-based models are evaluated (both by aboveground wood production and flux measurements). The second objective is to scale up these coupling models. It is important to assess forest production at regional and global scales. Myneni et al. (2001) have found that the wood biomass of northern forests is probably a main sink of the annual anthropogenic input of CO2. The regional scale also allows the parameterization and the evaluation of the models and is an interesting scale for evaluating global carbon models.
Few attempts have been made to spatially model and simulate forest growth processes at the regional scale. These models contain numerous parameters, making their application difficult at regional or larger scales because of the large amount of data required for model parameterization. Combined remote sensing and geographic information systems (GIS) tools provide new approaches for scaling up processes from the stand to regional or larger scales (Plummer 2000). Ditzer et al. (2000) presents such a study for a 55,000 ha tropical rain forest using the process-based forest growth model Formix 3-Q. That model, however, cannot account for the effects of climate change; photosynthesis is empirically based; and the model does not simulate soil water balance, soil carbon fluxes and heterotrophic respiration. Nevertheless, the model has proved useful for studies on allometry and forest management. We are interested in the Formix 3-Q modeling study because application of the model over the whole forest relies on the use of a GIS database that includes forest inventory data and aerial photographs.

In the present work, we used and parameterized a detailed process-based model to simulate, at a regional scale, the productivity and carbon budget of a large, managed temperate forest, characterized by high heterogeneity in species composition, tree age, canopy structure, soil type and water and mineral resources. Our approach was to use local measurements (large stand-scale database) and simulations performed stand by stand over the whole forest with a process-based ecosystem model. A distinctive feature of the study was that it linked forest inventories, soil inventories, satellite data and field measurements to produce a precise parameterization of the model and thus a precise quantification of forest carbon fluxes and productivity. This work is an essential first step to predicting both inter-annual and long-term climate change effects on forest carbon balance at the regional scale.

Materials and methods

Model description

CASTANEA is a physiological multilayer process-based model designed to predict the carbon balance of an even-aged, monospecific deciduous forest stand. The main output variables are: (1) leaf area index (LAI), standing biomass, soil carbon content and water content, which are state variables; and (2) canopy photosynthesis, maintenance respiration, growth of organs, growth respiration, soil heterotrophic respiration, transpiration and evapotranspiration, which are flux density variables (Davi 2004, Dufrêne et al. 2005).

The canopy is assumed to be horizontally homogeneous and vertically subdivided into a variable number of layers, each enclosing the same amount of leaf area (typically less than 0.2 m² m⁻²). No variability between trees is assumed and so one “average” tree is considered representative of the stand. A discussion of this assumption, based on an uncertainty analysis, is given in Dufrêne et al. (2005). The species used for the simulations are the main species of the stand, which may lead to inaccuracies in some cases, especially when mixed coniferous–deciduous stands are being considered. Tree structure is represented as a combination of foliage, stems, branches, coarse and fine roots. A carbohydrate storage compartment is also represented but not physically located in the model.

From incident radiation and photosynthetic characteristics of individual leaves, half-hourly rates of gross canopy assimilation and transpiration are calculated. Leaf nitrogen per unit area (N_L; g N m⁻² leaf) is calculated from measured leaf nitrogen concentration (N_m; g N g⁻¹ DM leaf), which is assumed to be constant inside the canopy, multiplied by leaf mass per area (LMA), which decreases exponentially inside the canopy. Photosynthetic capacity of leaves in different canopy positions is derived from leaf N_m.

After subtracting maintenance respiration requirements, the remaining assimilates are allocated to the growth of various plant tissues based on priorities that vary with season. The allocation coefficients depend only on the species. Fine roots and storage compartments were estimated by inverting the model assuming a constant biomass (i.e., equilibrium) on a yearly basis. This calibration was carried out at three sites, two in eastern France and one in the Netherlands: in Hesse for beech (Barbaroux 2002, Dufrene et al. 2005), in Champenoux for oak (Barbaroux 2002) and in Loobos for Scots pine (Davi 2004). The allocation coefficient for coarse roots was calculated assuming a constant ratio between coarse roots and trunks (coarse root:shoot ratio): a value of 0.2 was deduced from Cairns et al. (1997). Carbon allocation to aboveground wood is thus the resultant and is not calibrated, allowing us to use it as a validation output. An evaluation of this calibration in the Fontainebleau forest is given by Barbaroux (2002), and possible age and fertility effects are discussed below. Phenological stages (e.g., budburst, end of leaf growth, start of leaf yellowing) and leaf growth depend on degree days. Maintenance respiration depends on temperature and N_m of various organs, whereas growth respiration depends on the biochemical composition of organs.

The soil water balance sub-model includes three soil layers. The soil organic carbon sub-model (based on the Century model (Parton et al. 1987) and described in Epron et al. (2001)) separates soil organic carbon into three major components, which include active pools (live soil microbes plus microbial products), slow pools (resistant plant material) and passive pools (soil stabilized plant and microbial material) pools. Carbon flow between pools is controlled by decomposition rate and a microbial respiration loss parameter, which depend on soil texture, soil water and temperature. A major uncertainty in the soil model is the initialization of the carbon pools. We based the initialization on an equilibrium hypothesis.

There are two main time steps in the model: half-hourly and daily. The simulation period typically ranges from days to years. Most variables including fluxes (light penetration, photosynthesis, respiration, transpiration, rainfall interception, soil evaporation) are simulated half-hourly, whereas all the state variables (organ biomass, soil and carbon water content) and some others (growth and phenology) are simulated daily. The model includes several species sub-models, accounting
for the main species of France: Fagus sylvatica L., Quercus robur L./Quercus petraea (Matt.) Liebl. and Pinus sylvestris L. Other species sub-models may also be included (e.g., Quercus pubescens Willd., Quercus ilex L. and Pinus pinaster Ait.).

Input meteorological driving variables, either half-hourly or daily values, are global radiation, rainfall, wind speed, air humidity, temperature and atmospheric CO2 concentration. A list of simulated processes and the original authors of the sub-models are given in Table 1. A more complete description of the model, including equations, is given in Dufrene et al. (2005).

Most of the sub-models were tested and validated for F. sylvatica with independent measurements of canopy photosynthesis, autotrophic and heterotrophic respiration, NEE, wood and root growth, transpiration, soil water evaporation, rainfall interception and soil water status (Davi 2004, Davi et al. 2005). These measurements and validations were made at the Carboeuroflux Hesse flux-tower site (Granier et al. 2000) and are described in Davi et al. (2005). For the pine species, NEE, transpiration and evaporation fluxes were validated at the Loobos (Dolman 2002) and Le Bray Carboeuroflux sites (Berbigier et al. 2001). For the oak and beech species, model validation was made on annual growth calculated from wood drill cores extracted at 22 stands in the Fontainebleau forest (13 oak stands, nine beech stands) (Barbaroux 2002, Davi 2004).

The CASTANEA model has about 200 so-called parameters, most of which are constants: 15% are physical constants or regional parameters, and 65% are local-scale physical or biophysical constants describing the soil or the vegetation. These local-scale constants are dependent on tree species, but do not vary much from stand to stand (e.g., quantum yield). Finally, 20% of the parameters are variable from one stand to another, even if the species is the same (e.g., stand age and LAI). In our study, most of these spatially variable parameters were determined or estimated.

The sensitivity of the model to the parameters has been studied on a yearly basis with a Monte Carlo technique (Dufrene et al. 2005). Results have shown that the most sensitive parameters are the photosynthetic parameters, soil water-holding capacity parameters, canopy structure parameters and phenology driving parameters. Parameters that are both sensitive and highly spatially variable include LAI, aboveground woody biomass (AWB), age, $N_m$, soil depth and available soil water capacity.

**Study site**

The Fontainebleau forest, located in the southeast of Paris (48°25' N, 2°40' E) extends over 17,000 ha (Figure 1), with a mean altitude of 120 m. The region is characterized by a temperate climate, with a mean annual temperature of 10.6 °C and a mean annual precipitation of 750 mm fairly well distributed throughout the year. This forest is actively managed by the French National Forest Office (ONF) and is divided into about 3600 management units. Regular forestry practices modify the structure and species composition of the forest stands. The forest is composed of 38% oak ($Q. robur/Q. petraea$), 31% Scots
pine (P. sylvestris) and 11% beech (F. sylvatica). The remaining 20% are other evergreen or deciduous species (5%) and sparsely wooded area (15%). These percentages are calculated with the main species of the stands, but only 10% of the forest is mixed deciduous–coniferous stands. Understory species are hornbeam (Carpinus betulus L.) and birch (Betula pendula Roth.). Most of the commercially exploited stands are located in areas of flat topography (plain and plateau), while slopes of the forest are mostly rocky and sparsely wooded. The soil is mainly sandy because of Stampian sand parent rock and deposition of windborne sands. These windborne sands are composed of Stampian sand mixed with loam and clay, and cover the whole forest to different depths (Robin 1993). The range in soil type varies widely, going from rendzine to podzol with various intermediate brunisols. Some parts of the forest also present hydromorphy, leading to gleyic soils.

The forest shows a great variety of stands, including the successive stages of development: seedlings, thickets, sapling stands, pole stands, mature forest and seed tree stands. The range of species and soil types adds variability, so that the Fontainebleau forest captures major characteristics of a west paleoartic managed forest.

Available data set description

**Forest inventory attributes** Data are in a Geographical Information System built by the French National Forest Office in 1995. This GIS database contains a vectorized map of the boundaries of the forest stands based on aerial photography interpretation and field surveys. The resulting delimited stands are even-aged and homogeneous in species, structure, tree density, spatial distribution of trees and silvicultural practices. The Fontainebleau forest is divided into 3635 stands, with a mean area of 5 ha. Figure 1 shows the stand boundaries with the main species. For each stand, the attributes determined by measurements and site surveys are: (1) stand structure describing the distribution of trees by age and size classes; (2) basal area of the stand estimated at a tree height of 1.30 m with a basal area relascope at three locations in the stand; (3) dominant tree height, estimated by eye; (4) the three main species reported in order of cover importance—their mean diameter class and percent cover over the stand estimated visually; (5) the age class of the three main species estimated visually, stump analysis, archives references or wood drill core extraction, and the class intervals are 20 years for deciduous trees and 10 years for coniferous trees; and (6) other measured characteristics that were not used in our study.

**Soil database** The soil database was built by the French National Forest Office in 1995. It consists of about 8800 drilled pits evenly distributed over the forest (one every 2 ha). These measurements were done on a predefined spatial grid. Samples were obtained with a 2-m drill, and each description was made in the field. The attributes determined for each sample were soil type, humus type, underlying parent material and its depth if available, and for each horizon, the type, depth, texture, effervescence and pH. The nomenclature of soil and humus type was taken from Duchaufour (1982). The precision of the horizon thickness is less than 5 cm, and the total depth of the soil was reported when the drill blocked on the underlying parent material. Soil texture was determined as described by Baize (1988).

**Remote sensing data** On July 10, 1995, SPOT (Satellite Pour l’Observation de la Terre) satellite images of the region were acquired in three spectral bands: green, red (R) and near infra-red (NIR) and with a pixel size of 20 m. Images were rectified and geo-referenced using ground control points and integrated into the GIS database of Fontainebleau forest. Digital counts (gray tone) were converted to at-satellite (top of atmosphere)
radiance (W m\(^{-2}\) sr\(^{-1}\) \(\mu\)m\(^{-1}\)) using the gains and the offsets contained in the image headers and then calibrated to scaled surface reflectance after atmospheric corrections using a Dark Object Subtraction approach (Song et al. 2001). The classical Normalized Difference Vegetation Index (NDVI) is calculated for each pixel of the image as: \(\text{NDVI} = (\text{NIR} - R)/(\text{NIR} + R)\).

**Field measurements** Field measurements were made between 1994 and 2002 on 56 stands within the forest. The sampled stands represent the main species and a wide spectrum of ages, structures, soil compositions and tree densities. Most of the measurements in 1994 and 1995 were made during the European Multisensor Airborn Campaign (EMAC) organized by the European Space Agency (ESA), Ispra, Italy (Dufrêne et al. 1997). The measured variables were stand dendrometric characteristics: tree density, basal area, tree height, crown height, total biomass, trunk biomass and branch biomass, calculated as described by Proisy et al. (2000). In situ measurements of LAI were made with a Li-Cor LAI-2000 Plant Canopy Analyzer, from the end of June to mid-July 1995 under uniform diffuse sky on 41 of the 56 experimental stands. A detailed description and analysis of the spatial and temporal variability of LAI in the Fontainebleau Forest is given in Le Dantec et al. (2000). For leaf biochemical characteristics, sun and shade leaves were sampled to estimate leaf chlorophyll concentration, nitrogen concentration, LMA and leaf water content as described by Demarez et al. (1999). Stand age and annual wood increment of the stand were taken as the medians of the age and height classes for the main species.

**Species characteristics** Species, age and height were extracted directly from the forest inventory. For each stand, the species considered is the main species of the stand as registered in the Fontainebleau forest GIS database (the case of mixed stands is discussed later). The age and the height of each stand were taken as the medians of the age and height classes for the main species.

**Organ biomass and carbohydrate reserves** Branch, trunk, coarse root and carbohydrate storage biomass were derived from the total aboveground biomass of the stand (Figure 2 and Table 2). Aboveground woody biomass (g\(_{\text{DM}}\) ha\(^{-1}\) or g C m\(^{-2}\) based on the conversion factor 0.5) was estimated for the EMAC stands from field measurements and allometric equations as described by Proisy et al. (2000). The product of basal area (BA) and dominant height (H) is linearly related to AWB. Two relationships were calibrated for deciduous and evergreen species (Le Dantec 2000):

\[
\text{AWB} = aBAH + b
\]

where \(a = 0.37\) or 0.22 and \(b = -0.01\) or 4.6, respectively, for deciduous \((r^2 = 0.96, 35\) stands) or coniferous stands \((r^2 = 0.99, 13\) stands).

**Leaf area index, leaf biomass and fine root biomass** Leaf area index was estimated based on within-stand NDVI variability (Davi 2004). For each stand, LAI was calculated based

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### Table 2. Initialization of biomass compartments of the stands. Input data are leaf area index (LAI), aboveground wood biomass (AWB) and leaf mass per area (LMA). Abbreviation: LWB = living wood biomass.

<table>
<thead>
<tr>
<th>Initialization</th>
<th>Quercus petraea/Q. robur</th>
<th>Fagus sylvatica</th>
<th>Pinus sylvestris</th>
</tr>
</thead>
<tbody>
<tr>
<td>Aboveground wood biomass (AWB)</td>
<td>Input data</td>
<td>Input data</td>
<td>Input data</td>
</tr>
<tr>
<td>Branch biomass(^1)</td>
<td>If age &gt; 175, 0.37AWB</td>
<td>If age &gt; 175, 0.37AWB</td>
<td>0.13AWB</td>
</tr>
<tr>
<td></td>
<td>If age &lt; 25, 0.13AWB</td>
<td>If age &lt; 25, 0.13AWB</td>
<td></td>
</tr>
<tr>
<td></td>
<td>Else (0.0016age + 0.1)(AWB)</td>
<td>Else (0.0016age + 0.1)(AWB)</td>
<td></td>
</tr>
<tr>
<td>Living wood biomass (LWB)(^2)</td>
<td>1.3age(^{-0.52})AWB</td>
<td>4.3age(^{-0.87})(AWB)</td>
<td>0.9age(^{-0.52})AWB</td>
</tr>
<tr>
<td>Leaf biomass (max.)</td>
<td>= Leaf biomass (max)</td>
<td>= Leaf biomass (max)</td>
<td>= Leaf biomass (max)</td>
</tr>
<tr>
<td>Fine root biomass(^3)</td>
<td>0.2AWB</td>
<td>0.2AWB</td>
<td>0.2AWB</td>
</tr>
<tr>
<td>Coarse root biomass(^4)</td>
<td>0.3LWB</td>
<td>0.3LWB</td>
<td>0.25LWB</td>
</tr>
<tr>
<td>Carbohydrate storage biomass(^5)</td>
<td>0.3LWB</td>
<td>0.3LWB</td>
<td>0.25LWB</td>
</tr>
</tbody>
</table>

\(^1\) Hatsch (1997).  
\(^2\) Ceschia et al. (2002).  
\(^4\) Cairns et al. (1997).  
\(^5\) Barbaroux et al. (2003).
on an empirical relationship between ground-based LAI measurements and SPOT reflectance data calibrated over 40 stands (12 dominated by beech, 17 by oak and 11 by Scots pine) (Figure 3).

The SPOT satellite image was acquired in July 1995, which coincides with the period of maximum LAI of the stands (the model needs the maximal leaf area of the year and computes the increase and decrease in LAI in spring and autumn). After making the geometrical and atmospheric correction procedures outlined above, NDVI was calculated for each image pixel. Then, based on the forest polygon coverage of forest stands, the standard deviation of NDVI was calculated for each stand. To avoid edge effects and overlapping pixels between stands, pixels located in a 20-m wide band along the polygon limits inside the stand were excluded. Stands composed of less than 25 pixels (1 ha) were not considered. For stands including more than 25 pixels, LAI was calculated as:

\[
\text{LAI} = -3.59 \log(\text{std(NDVI)}) - 11.22
\] (2)

where std(NDVI) is the standard deviation of the NDVI inside the stand \((r^2 = 0.87, n = 40)\).

The remotely sensed LAI obtained from Equation 2 is underestimated in coniferous stands because of clumping of needles, branches and trees. To correct for clumping, the remotely sensed LAI of the coniferous stands was divided by the clumping factor of 0.57 reported by Stenberg et al. (1994) for Scots pine.

Total leaf biomass of the stand was calculated from LAI and the LMA profile integrated over the canopy. Fine root biomass was assumed equal to leaf biomass (Table 2) (Vogt et al. 1987, Santantonio 1989, Bauhus and Bartsch 1996).

Spatialization of leaf nitrogen concentration \((N_m)\) over 3600 forest stands is based on two relationships between soil type and leaf \(N_m\) measured in 40 stands of the Fontainebleau forest, plus one measurement from Brèthes and Ulrich (1997) for the podzol soil (Figure 4). Nine soil types are present in the Fontainebleau forest. They are clas-

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**Figure 2.** Relationship between aboveground woody dry biomass (MgDM ha\(^{-1}\)) and the product of basal area (m\(^2\) ha\(^{-1}\)) and dominant height (m), based on data from 48 stands in the Fontainebleau forest. Regression equations are given in the text (Equation 1) (from Le Dantec 2000).

**Figure 3.** Relationship between mean stand leaf area index measured in the field and the standard deviation of the Normalized Difference Vegetation Index (NDVI) of the SPOT satellite pixels that are inside the stand limits. Regression equation is given in the text (Equation 2) (from Davi et al., unpublished results).

**Figure 4.** Relationship between mean leaf nitrogen concentration (mg g\(^{-1}\)) and mean stand soil type. Soil types are indicated by numbers 1 to 9 (the “brown podzol above luvisol” corresponds to two superposed soils—a new soil (brown podzol) developed on windborne sands, and an older luvisol).
sified according to their fertility from 1 (rendzina) to 9 (podzol), with intermediate types of brunisol. To spatialize $N_m$, determination of the mean stand soil type is necessary, although difficult, either because many soil types may be found in the same stand or because there was no soil pit within particular stands. These difficulties were resolved by soil type interpolation. We used the linear inverse distance weighted interpolation method between adjacent points. The method was first tested by setting aside 20% of the soil types and by predicting both the soil type value of the removed pits and the soil type value of the stand that contained these pits. The standard deviation of the predicted difference is 1.28 in the case of removed pits and 0.38 for soil type predictions at the stand scale. A nearest neighbor assignment method gave poorer results (1.62 and 0.46, respectively). A new soil type map was then compiled with a pixel of 5 m, with soil value being a real number. This method accounts for the spatial representativity of each soil pit, and attributes a soil type for even the smallest stands. The stand soil type is the mean value of every pixel contained in the stand, leading to real values of stand soil type. The standard deviation of these pixel values is an estimate of the within-stand spatial heterogeneity of soil. Highly heterogeneous stands (standard deviation greater than one, 101 stands) were not taken into account and a mean $N_m$ value was assigned to these stands based on measurements from the Fontainebleau forest experimental stands (22.8 mg g$^{-1}$ for oak, 24.4 mg g$^{-1}$ for beech and 14.7 mg g$^{-1}$ for pine).

This procedure was applied to the experimental stands and two reliable relationships with leaf nitrogen were found, one for deciduous stands and another for evergreen stands (Figure 4). These relationships were applied to all other stands of the forest to estimate $N_m$ based on the “mean” soil type of the stand.

**Soil depth**  
Soil depth, which is the depth between the soil surface (without humus) and the underlying parent material, is used for the calculation of available soil water capacity. The long and complex geological history of the Parisian basin accounts for the highly variable soil depth. For some drills, soil depth was derived directly from the database, for example, if the drill was obstructed by a sandstone or limestone flag. But for other locations, the soil depth was not obvious, either because the drill was obstructed before reaching the underlying material or because the underlying material was too deep (more than 2 m). Several hypotheses were formulated to calculate soil depth when it was not directly available from the database. Where the drill was obstructed before reaching the underlying material, we added up to 100 cm depending on the soil type, the effervescence, the texture and the soil depth from other pits surrounding the soil pit. Soil depth was interpolated between adjacent soil pits, with the same procedure as for soil type. We then calculated mean soil depth for each stand.

**Sand and clay percent of the top and deep soil**  
The texture class of each horizon at each soil pit was converted to percent of silt, clay and sand, based on the soil texture triangle (Baize 1988), taking the central value for each texture category. For each pit, the percentages of silt, sand and clay were determined for the top soil (0–30 cm) and for the deep soil (30 cm to underlying parent material depth) and weighted by horizon thickness. These values were interpolated by the procedure described for soil type. We then calculated these parameters for every stand of the forest.

**Soil water characteristics**  
Wilting point, field capacity and available soil water capacity for the top and total soil were determined with pedotransfer functions, where inputs are the soil percent of sand and clay (Saxton et al. 1986). These parameters were calculated by horizon, then multiplied by horizon thickness and summed to give top soil values (0–30 cm) and total soil values (0 cm to underlying material depth). The soil is mainly sandy and generally does not contain stone. For each pit, we obtained a wilting point, field capacity and available soil water capacity for the top and for the total soil. These values were interpolated using the procedure described for soil type, and mean stand values were calculated.

**Soil carbon content initialization**  
Initial soil carbon content was determined for each stand by simulation based on the equilibrium hypothesis (the annual carbon balance of the soil is zero). The annual equilibrium for the different carbon pools was calculated by a direct numerical resolution technique for the year 1990. Then 4 years of CASTANEA simulations were run without reinitializing the soil carbon pools between years. These 4 years of runs increase the reliability of the simulations by taking into account inter-annual variability of climate, in particular for the carbon pools with high turnover. With this method, we obtained the initialization of all the carbon pools for the year 1995. Numerical resolution of the equilibrium is a fast method, allowing us to avoid the time-consuming 200 years of climate time-series simulations that is commonly used to find the equilibrium. The method showed good similarity with 200 years of simulations (1960–2000 climatic data repeated five times) for soil respiration and total carbon content (not shown). With our methodology, the soil carbon is not quite in equilibrium for the year 1995.

**Model run specifications**  
Rocky zones and sparsely wooded stands with superficial soils were excluded from the simulations (643 stands removed): their wood production and carbon fluxes are considered negligible. Other stands (2992) were parameterized with their main species submodel (for *Q. robur*/*Q. petraea*, *P. sylvestris* or *F. sylvatica*) and with the parameters that were spatially determined. For the other species (5% of the studied area), we used biochemical and biophysical parameter values from species that are ecologically comparable. Parameters that were not spatially determined were set equal to the mean value measured on the experimental stands for each species. Some parameters were set to their mean value by default because no spatialization was possible (e.g., LMA). When no measurements were available in Fontainebleau, parameters were set to values from other French sites (Le Bray, Hesse) or from the literature. The run was performed on each stand, assuming no interaction between stands. The climate was not spatialized over the forest. Initial climatic data were taken from daily time se-
ries recorded at one meteorological station located near the forest.

Cumulative annual fluxes are the main outputs: gross primary production (GPP) is total photosynthetically fixed carbon. Autotrophic respiration (Ra) is the sum of growth and maintenance respiration of the tree organs. Net primary productivity (NPP) is the difference between the GPP and Ra; i.e., it includes the carbon stored in the vegetation during the year plus the carbon in biomass that died in that year. Aboveground woody increment (AWI) is the annual growth of the aboveground woody trunk (trunk + branches). Heterotrophic respiration (Rh) is the respiration of soil organisms (mainly microbial) leading to soil carbon efflux. Net ecosystem exchange (NEE) is the difference between NPP and Rh, and is a measure of the net carbon fluxes between the forest ecosystem and the atmosphere. This parameter determines if the forest ecosystem is a source or a sink of carbon and to what degree. The sign convention chosen is the ecosystem-based convention: GPP is positive, respiration is negative, NEE is either negative (carbon source) or positive (sink).

The ONF and IFN provide estimates of the Fontainebleau forest commercial woody volume and annual volume increase. To compare our results with these estimates for the year 1995, we have taken only the trunk into account: the trunk percentages used for the calculations are the same as in Table 2. Trunk carbon biomass and trunk biomass increase were converted to m$^3$ or m$^3$ year$^{-1}$ based on percent of carbon in dry matter (0.5 g C/g DM) and wood density (570 kgDM m$^{-3}$ for Quercus petraea, 550 kgDM m$^{-3}$ for Fagus sylvatica and 460 kgDM m$^{-3}$ for Pinus sylvestris). Finally, trunk volume was converted to large roundwood volume with a biomass expansion factor function of the stand age (J.-F. Dhoit, INRA, Nancy, France, personal communication, for deciduous spp., and Lehtonen et al. (2004) for Pinus). The volume and volume increment of each stand were summed to obtain the total volume and total volume increment of the forest. The percent increase was calculated as the ratio of these values.

Results

Parameterization results

The model input parameters attributed to each stand are means and are reported as area-weighted for the three main tree species and for the entire forest (Table 3). The histograms presented in Figure 5 represent the cumulated area of each parameter class. Stand LAI values range from less than 1 to 8. Low values represent seed tree stands or sparsely wooded areas; high values are dense closed canopies. The leaf $N_o$ histogram is bimodal: the low values represent coniferous stands, and the high values represent deciduous stands. The distribution for coniferous stands is less spread out, showing that coniferous stands are mainly on the same type of soil (brown podzol), whereas the deciduous stands are mainly on argilic brown soils. Soil depth varies widely in the Fontainebleau forest, ranging from 100 to more than 200 cm. Aboveground woody biomass and age are typical for a managed forest comprising many different-aged stands. Available soil water capacity is low, mainly because of the sandy soils, but can reach values of 200 mm.

Simulation results

The simulation results are weighted by stand area and averaged over the whole forest (Table 4). Each output is expressed as g C m$^{-2}$ year$^{-1}$. Histograms are also presented (Figure 6). Annual GPP values range from low values (low LAI stands) to more than 2000 g C m$^{-2}$ year$^{-1}$ (Scots pines stands with high LAI). The annual NPP histogram shows some negative values, perhaps indicating errors in parameterization as a result of errors in the data set; e.g., overestimation of wood biomass of the stand would lead to overestimation of autotrophic respiration. The NEE is mainly positive but some stands show negative values (i.e., carbon source).

Maps of AWI and NEE are presented in Figure 7. They illustrate the high variability in stand functioning, creating a complex mosaic with apparently no spatial correlations. However, the species spatial distribution explains a part of the spatial distribution in NEE and AWI (Figure 1).

Trunk wood volume of the forest is about 2.6 × 10$^6$ m$^3$ and the volume increment is 7.1 × 10$^4$ m$^3$ (2.74%) for the year 1995 (Table 5). Estimates made by IFN and ONF are given for comparison, either for the entire forest or for the principal species.

Discussion

Parameterization methodology

Unlike statistical empirical models, the CASTANEA model is a process-based ecophysiological model that requires many

Table 3. Mean values of key parameters scaled stand by stand over the Fontainebleau forest. Results are weighted by stand area and averaged by species, and calculated for the whole forest (Total).

<table>
<thead>
<tr>
<th>Species</th>
<th>LAI</th>
<th>Leaf nitrogen concentration (mg g$^{-1}$)</th>
<th>Aerial wood biomass (g C m$^{-2}$)</th>
<th>Age (year)</th>
<th>Soil depth (m)</th>
<th>Available soil water capacity (mm)</th>
<th>Total soil carbon (g C m$^{-2}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Quercus petraea/Q. robur</td>
<td>3.1</td>
<td>22.8</td>
<td>10,980</td>
<td>137</td>
<td>154</td>
<td>133</td>
<td>4,774</td>
</tr>
<tr>
<td>Fagus sylvatica</td>
<td>3.1</td>
<td>23.4</td>
<td>8,767</td>
<td>89</td>
<td>133</td>
<td>121</td>
<td>3,850</td>
</tr>
<tr>
<td>Pinus sylvestris</td>
<td>2.1</td>
<td>14.9</td>
<td>5,851</td>
<td>68</td>
<td>157</td>
<td>124</td>
<td>5,277</td>
</tr>
<tr>
<td>Total</td>
<td>3.0</td>
<td>20.6</td>
<td>8,854</td>
<td>106</td>
<td>150</td>
<td>128</td>
<td>4,796</td>
</tr>
</tbody>
</table>
input parameters. We were able to obtain most of these parameters for the Fontainebleau forest. The CASTANEA model was created to simulate the carbon and water fluxes for homogeneous forest stands, and was validated on homogeneous stands with flux tower and growth data. However, the parameterization of the model at this scale proved difficult, and an unknown source of error or bias may have been introduced.

The stands are considered to be even-aged and homogeneous in species, structure, tree density and spatial distribution of trees. This is the case for most of them (it was the polygon delineation criterion). However, some stands may have an irregular structure, which may create bias in the results because of the non-linearity process. A similar bias may also occur when stand mean parameters values are used.

We assumed no interaction among stands, i.e., each stand functions independently of the surrounding stands. This hypothesis is justified for the hydrology because: (1) the topography is flat in the wooded areas; and (2) because the soil is sandy and deep, so there is essentially no horizontal runoff. Another possible interaction is shadowing, especially if a tall

Table 4. Carbon flux (g C m$^{-2}$ year$^{-1}$) outputs averaged over the Fontainebleau forest. Results are weighted by stand area and averaged by species, and calculated for the whole forest. Abbreviations: GPP = gross primary productivity; NPP = net primary productivity; and NEE = net ecosystem exchange.

<table>
<thead>
<tr>
<th>Species</th>
<th>GPP</th>
<th>Autotrophic respiration</th>
<th>NPP</th>
<th>Aboveground woody biomass increment</th>
<th>Heterotrophic respiration</th>
<th>NEE</th>
</tr>
</thead>
<tbody>
<tr>
<td><em>Quercus petraea/Q. robur</em></td>
<td>1408</td>
<td>–737</td>
<td>672</td>
<td>244</td>
<td>–263</td>
<td>409</td>
</tr>
<tr>
<td><em>Fagus sylvatica</em></td>
<td>1225</td>
<td>–693</td>
<td>532</td>
<td>252</td>
<td>–232</td>
<td>301</td>
</tr>
<tr>
<td><em>Pinus sylvestris</em></td>
<td>1357</td>
<td>–699</td>
<td>658</td>
<td>260</td>
<td>–268</td>
<td>390</td>
</tr>
<tr>
<td>Total</td>
<td>1355</td>
<td>–725</td>
<td>630</td>
<td>248</td>
<td>–260</td>
<td>371</td>
</tr>
</tbody>
</table>
Figure 6. Histograms of the distribution of output fluxes (g C m\(^{-2}\) year\(^{-1}\)): gross primary production (GPP), autotrophic respiration (Ra), net primary productivity (NPP), aboveground wood increment (AWI), heterotrophic respiration (Rh) and net ecosystem exchange (NEE). The x-axis is the parameter class and the y-axis is the cumulated area for the considered class.

Figure 7. Maps of annual aboveground wood increment and net ecosystem exchange simulated over the Fontainebleau forest in 1995.
leaf that soil fertility is accounted for in the simulations through
Schjoerring 2003). An advantage of the relationship we used is
and Bausch 2003, Graeff and Claupein 2003, Hansen and
mate of leaf includes a wide range of soil types, allowing for a rough esti-
plained on the basis of soil fertility. The Fontainebleau forest

<table>
<thead>
<tr>
<th>Species</th>
<th>IFN volume (m$^3$)</th>
<th>ONF volume (m$^3$)</th>
<th>Volume estimated (m$^3$)</th>
<th>IFN volume increment (m$^3$ year$^{-1}$)</th>
<th>Volume increment (m$^3$ year$^{-1}$)</th>
<th>IFN RGR (% vol. year$^{-1}$)</th>
<th>Estimated RGR (% vol. year$^{-1}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Quercus petraea Q. robur</td>
<td>1,297,281</td>
<td>1,076,303</td>
<td>1,344,000</td>
<td>23,923</td>
<td>28,000</td>
<td>1.84</td>
<td>2.08</td>
</tr>
<tr>
<td>Fagus sylvatica</td>
<td>414,390</td>
<td>333,589</td>
<td>325,520</td>
<td>14,254</td>
<td>8,940</td>
<td>3.44</td>
<td>2.75</td>
</tr>
<tr>
<td>Pinus sylvestris</td>
<td>732,796</td>
<td>655,267</td>
<td>637,150</td>
<td>21,989</td>
<td>25,123</td>
<td>3.00</td>
<td>3.94</td>
</tr>
<tr>
<td>Other species</td>
<td>266,057</td>
<td>227,527</td>
<td>295,400</td>
<td>8,008</td>
<td>9,106</td>
<td>3.01</td>
<td>3.08</td>
</tr>
<tr>
<td>Total</td>
<td>2,710,524</td>
<td>2,292,686</td>
<td>2,602,070</td>
<td>68,174</td>
<td>71,170</td>
<td>2.52</td>
<td>2.74</td>
</tr>
</tbody>
</table>

stand is located south of a small stand, but this effect is likely
to be small. Finally, an interaction could be more climatic, for
example wind speed or relative humidity, but these effects are
difficult to quantify and include in a model.

The climate was not spatialized. A fine grid of meteorologi-
cal data was computed with the model AURHELY developed
by Météo-France (Bénichou and Le Breton 1987). This model
uses meteorological data from local stations and interpolates
the data taking account of topography. A preliminary analysis of
these interpolated data showed only small differences with
the mean value in the areas of interest; therefore, the hypothe-
sis of a mean climate over the forest seems justified, but needs
further investigation.

Leaf area index was deduced from a relationship with
the standard deviation of the NDVIs of the stand. The use of poly-
gen-based aggregation of remotely sensed data for LAI esti-
cation can provide more information than pixel-based
 estimations (Wicks et al. 2002). The NDVI standard deviation
is a quantification of the heterogeneity inside the stand. In a
managed forest, the LAI of a stand is closely linked to hetero-
genecity: a small LAI implies open canopies, with large gaps
(e.g., seed tree stands), leading to high heterogeneity, whereas
a high LAI is obtained when the canopy is closed and more ho-

genous. A second explanation for the close correlation be-

 tween the NDVI standard deviation and ground-measured LAI
is the saturation of the LAI–NDVI relationship: at low LAI,
this relationship is quite linear, so for small changes in LAI
from one pixel to another inside the stand, there is a propor-
tional change in NDVI. But at high LAI, a small change in LAI
only slightly affects the pixel NDVI value, as reflected in the
higher NDVI standard deviation for low LAI distributions, and
the smaller NDVI standard deviation for high LAI distribu-
tions.

The relationship between soil type and leaf $N_m$ can be ex-
plained on the basis of soil fertility. The Fontainebleau forest
includes a wide range of soil types, allowing for a rough esti-
mate of leaf $N_m$. Work is being done to develop a more precise
estimate of $N_m$ based on hyperspectral satellite images (Diker
and Bausch 2003, Graeff and Claupein 2003, Hansen and
Schjoerring 2003). An advantage of the relationship we used is
that soil fertility is accounted for in the simulations through
leaf $N_m$, which is directly linked to leaf photosynthetic capa-
city. Stands located on poor soils (e.g., brown podzol) will
have lower leaf $N_m$, leading to lower photosynthesis and
growth compared with stands on fertile soils.

The soil equilibrium hypothesis assumes that soil carbon in-
puts (litter fall and root fall) approximate carbon outputs from
heterotrophic respiration. The difference between these two
fluxes depends on annual climate variations. This hypothesis
has previously been tested for soil respiration efflux and gives
good agreement for deciduous species (Epron et al. 2001,
Yang et al. 2002). For coniferous species, we think that the
simulated fluxes are overestimated: under a coniferous can-
oply, there is generally an accumulation of carbon (King 1995),
which is in contradiction with the equilibrium hypothesis.

Simulation of carbon fluxes over a large forest requires
great simplifications: many parameters are extremely respon-
sive in the model and highly spatially variable (e.g., LMA), but
are not easily measured at this scale. The spatialization of
LMA could be done through remote sensing, but no such LMA
data for a forest canopy are currently available (Blackburn and
Steele 1999).

We did not consider the effects of the mixing of species in-
side a stand, either involving the three main species (oak,
birch or birch). In our simulations, the species attribution for each
stand is the main species of the stand. For deciduous mixed
species, the effect on stand flux should be low because we esti-

mated LAI based on the total leaf area of the stand. The basal
area is also the total of the stand. Therefore the small variation
caused by the slightly differing functioning of oak and beech
would be simulated. For mixed coniferous–deciduous species,
use of the mean parameters of the dominant species is inade-
quate because phenology, photosynthesis and respiratory pro-
cesses differ greatly among species. However, because the
stands of mixed coniferous–deciduous species represent less
than 10% of the forest, if the criterion for heterogeneity is that
the dominant species represents less than 80% of the canopy
cover of the stand, we suppose that the errors thus made would
be negligible. Finally, the understory coverage can have a large
effect on stand water fluxes. The understory modifies stand
carbon fluxes directly (understory GPP and Ra) or indirectly
through soil water status. It is difficult to account for the
understory because we have little knowledge of its function in
terms of carbon and water fluxes and it is difficult to map it
over the whole forest.
The rocky areas, representing about 15% of the total forest surface, were not simulated. These areas, with superficial soils, are sparsely wooded, mostly boulder covered and sloping. Their wood production and carbon fluxes were considered to be negligible compared with the rest of the forest. For the forest area with species for which the CASTANEA model was not parameterized, which represent only about 5% of the forest area, the carbon fluxes were simulated with parameters from ecologically comparable species. Comparison with the ONF and IFN estimates of wood volume and annual volume increase was realized separately for the three dominant species.

Simulations

For deciduous species, Fluxnet towers measurements of GPP were between 1100 and 1700 g C m⁻² year⁻¹ and between 700 and 2200 g C m⁻² year⁻¹ for coniferous species, depending on location, year and stand age. For deciduous stands in a temperate climate, GPP measurements range from 950 to 1700 g C m⁻² year⁻¹ (4 sites: Hesse, France; Harvard Forest, MA, USA; Walker Branch, TN, USA; Willow Creek, WI, USA). For coniferous stands in a temperate climate, GPP measurements range between 1100 and 1500 g C m⁻² year⁻¹ (3 sites: Duke, NC, USA; Loobos, Netherlands; and Metolius, OR, USA) (Law et al. 2002, Valentini et al. 2000). Thus, our GPP simulations fall within the range of measured GPP values.

Measurements of NEE at these same Fluxnet sites ranged from −50 to 870 g C m⁻² year⁻¹ for deciduous species and from 260 to 650 g C m⁻² year⁻¹ for coniferous species. The values obtained in our simulations fall into these ranges for deciduous and evergreen stands on average, and for most of the stands separately (Figures 6 and 7).

Our simulated heterotrophic respiration rates are questionable, especially the finding that respiration in coniferous stands is of the same order of magnitude as in deciduous stands. The simulations may indicate that the approach used to initialize the carbon pools is inadequate or that there is little litter accumulation (equilibrium hypothesis). This result directly affects simulated NEE.

Our simulated NPP in Fontainebleau (630 g C m⁻² year⁻¹) is only slightly higher than the MODIS value of 590 g C m⁻² year⁻¹ derived from large-scale satellite based calculations (Turner et al. 2003). We compared our results with measurements made by both IFN and ONF for the whole Fontainebleau forest (Table 5). Our calculation of volume is in agreement with the IFN estimates for the whole forest, with a 4% difference in the total; however, there is a 14% difference with the ONF volume estimates.

Simulated volume increment is also in good agreement with the IFN estimation (5% difference in total). Trunk relative growth ratio (RGR, volume increment divided by the volume and expressed as a percentage increment per year) provides a way to compare our forest productivity estimates with those of the IFN. Our RGR values are in good agreement with the IFN values for the Quercus species, slightly lower for Fagus and overestimated for Pinus. In total, we overestimated forest productivity by 9%.

Possible reasons for these discrepancies are: (1) the year 1995 had favorable climatic conditions, with more precipitation than normal, limiting the autumn drought, so NPP was about 10% higher than the mean over 30 years of climatic data; (2) the conversion from daily to half-hourly climatic data leads to an overestimation of NPP of about 10%, because days with high radiation variability were not simulated in the half-hourly transformed data; and (3) the possible presence of nutrient limitation, especially phosphorus (P), which is unaccounted for in the model. Mean leaf phosphorus concentration in the Fontainebleau forest is around 1 mg g⁻¹ for beeches and oaks and 0.9 mg g⁻¹ for Scots pines. For comparison, mean P values of various forests in France are generally 1.35 mg g⁻¹ for oaks, 1.24 mg g⁻¹ for beeches and 1.33 mg g⁻¹ for Scots pines (Croisé et al. 1999). Phosphorus limitation can affect forest productivity (Bauer et al. 1997, Loustau et al. 1999, Warren and Adams 2002).

The model may overestimate or underestimate aboveground wood increment because of the simplicity of the carbon allocation module. This module ignores possible variations between stands differing in age and fertility. In aged and nutrient-poor stands, carbon allocation to fine roots could be enhanced (Magnani et al. 2000) and therefore greater than the constant value used in the model. Moreover, the older the stand, the greater its biomass and simulated autotrophic respiration, which leads to a greater use of carbohydrate storage during the winter. So, allocation to reserves, which was calibrated for young stands, could be underestimated for the oldest stands. If these two carbon sinks were underestimated for part of the forest, we probably overestimated the aboveground wood increment. Nevertheless, a bias in the carbon balance between aboveground and belowground compartments will have little effect on the NPP and NEE estimates because fine roots and aboveground wood have similar construction and maintenance costs.

Conclusion

We attempted to estimate carbon fluxes of a whole forest. Our results underline the importance of precise parameterization and validation. Our approach, consisting of stand-by-stand simulations, with attributes derived from forest inventories, soil inventories, satellite data and field measurements, shows good potential. More field measurements are needed to spatially validate our results. Soil respiration measurements are now in progress, as well as growth estimates from stem cores. We used an ancillary growth data set to improve the accuracy of the carbon allocation model, particularly to reproduce the decline in NPP with stand age more accurately (Gower et al. 1996). Our productivity results differ slightly from, but are close to, the growth estimates of the IFN. Reasonably good agreement was found between simulated trunk RGR (2.74%) and regional production estimates by the IFN (2.52%), as well as between simulated and measured annual wood production at the forest scale (around 71,000 m³ year⁻¹ and 68,000 m³ year⁻¹, respectively). This work contributes to the effort to
model climate change effects on forest productivity at the regional scale.

Acknowledgments

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